

Alphonsus crater with 11 pyroclastic vents

# Rationale for Landing Sites at Lunar Pyroclastic Deposits

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- Overview of the lunar pyroclastic deposits
- Resource potential
- ISRU, Accessibility and Traversability
- Recommendations for lunar landing sites

# **Explosively emplaced volcanic deposits**

Diffuse boundaries, association with vents

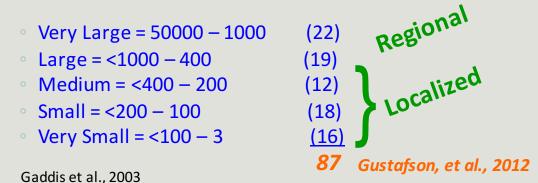
## Ancient, by association with lunar maria

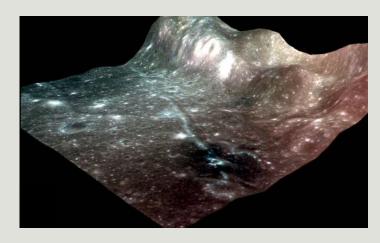
- Most basalts erupted during the Late Imbrian (3.6 to 3.8 BY ago)
- Hiesinger et al. (2000) crater counts

## Globally distributed

- Observed along margins of lunar maria
- Often associated with floor-fractured craters in faulted regions

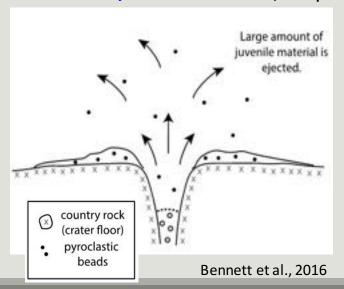
# Range of sizes (~50,000 to 3 km<sup>2</sup>):

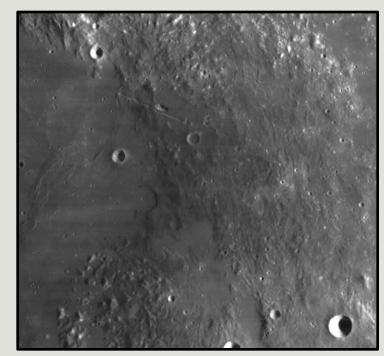




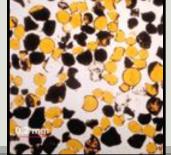
Alphonsus NE floor vents Kaguya MI color, 20 m/pixel, ~24 km across

- Formed by fire-fountaining (Head & Wilson, 2017)
  - High volatile content (H, H<sub>2</sub>O, CO, SO<sub>2</sub>), ballistic trajectories
  - Widely distributed, ~thin deposits
  - Abundant juvenile material
    - Quenched glass, crystalline beads (often Hi TiO<sub>2</sub>)
- Type Examples
  - Taurus Littrow (Apollo 17)
  - Rima Bode, Sinus Aestuum, Sulpicius Gallus





Rima Bode Regional Pyroclastic Deposit (~6600 km²) Lunar Orbiter IV 109, ~120 km across



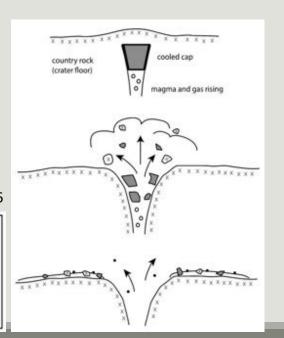
Ap 17 glass & crystalline beads (G. Ryder)

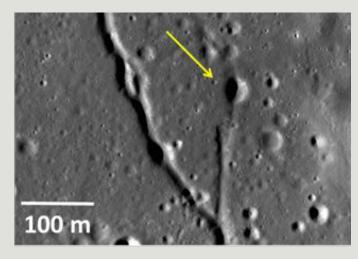
#### Formed by Vulcanian-style eruption

- Intermittent, violent explosion caused by degassing near surface, disruption of a plug as magma rises in a dike (Head & Wilson, 1979)
- **Volatile:** CO-rich gas produced by graphite oxidation (*Fogel & Rutherford, 1995*)
- Mixed juvenile and non-juvenile materials
  - Glass, crystalline basaltic (Hi FeO), fragmented country rock

#### Type Example

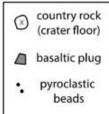
- Alphonsus crater
- Oppenheimer crater

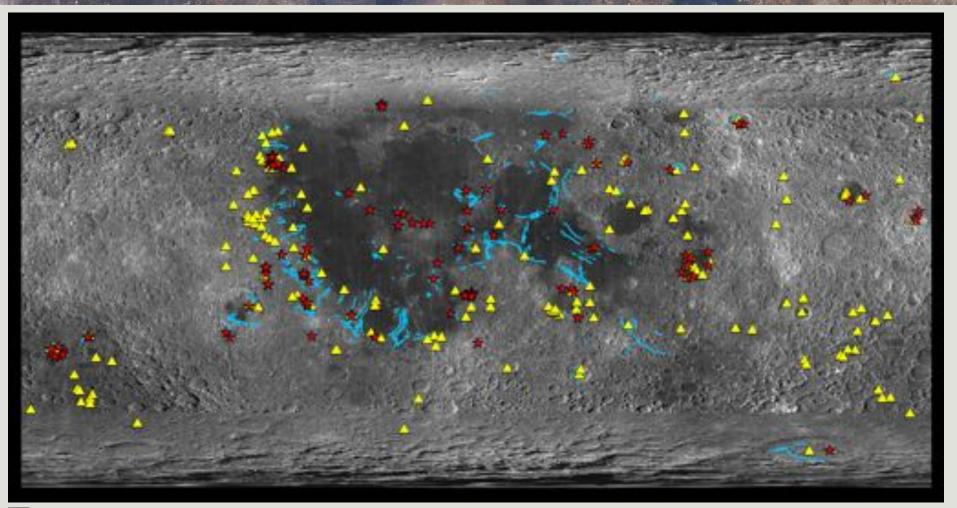


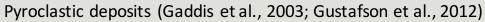


Alphonsus crater, Floor Vent #6 Kaguya Terrain Camera

Bennett et al., 2016







Floor-fractured craters (Jozwiak et al., 2015)

Fractures/faults (Wilhelms, 1987)

#### Comprised of 'primitive' glass and quenched spheres

- Relatively unprocessed materials from deep within the lunar mantle (~300 to 400 km depths)
  - High Mg/Al, MgO, mg#; lower Al<sub>2</sub>O<sub>3</sub> and CaO than mare basalts
- Interior crystals of olivine, spinel, ilmenite needles, etc.
- Uniform grain size, ~40 microns
- Surficial geochemical enrichments in >25 volatile elements
  - e.g., Au, Ag, Cu, Cd, F, S, Z (McCubbin et al., 2015)
- Contain FeO: 16 to 24 wt %
- Variable amounts of TiO<sub>2</sub>, related to color

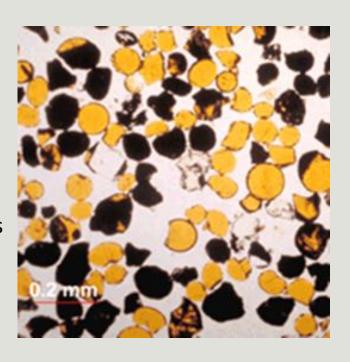
Green: <1 to 5 wt%</li>

Yellow: 5 to 9 wt %

Orange: 9 to 14 wt %

Red-black: >14 wt%

 25 varieties of volcanic glass described by Delano (1986), several others identified since



Ap 17 spheres viewed through a microscope. These have ~8 wt% TiO2. (G. Ryder)

#### **Indigenous Magmatic Water**

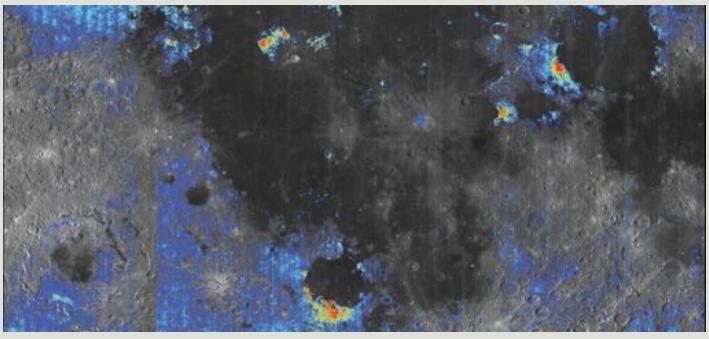
- The Moon is generally volatile-depleted
- •BUT, recent sample analyses of melt inclusions in glass samples from Ap15, Ap17 found evidence for water and other magmatic volatile species
  - ~30 to 36 ppm of magmatic water in Ap 15 glass beads (Saal et al., 2008)
  - Correlations with other magmatic elements (Chlorine, Fluoride, Sulfur)
  - Concentration of volatiles decreased toward edges, indicating they were not contaminants from Earth
  - 270-1200 ppm H<sub>2</sub>O in olivine crystals from Ap17 (Hauri et al., 2011)
    - From a primary lunar magma
    - Source water abundances of 80-440 ppm, comparable to MORBs
  - Summarized by McCubbin et al., 2015; Hauri e al., 2017



Apollo 15 green glass beads about 0.1 mm across (Ap 15 S79-32188)

#### **Mapping of Lunar Water**

- Use of Moon Mineralogy Mapper (M3) data, quantitative analysis of strength of the 2.85-micron band (ESPAT), mapping of OH- and/or  $H_2O$ -bearing minerals
- Water abundances up to 150 ppm at large pyroclastic deposits, 300 to 400 ppm near vents



ESPAT=effective single particle absorption thickness

ESPAT	,≤0.001 •	0.02	0.03	0.04	0.05	≥0.06
.wt% H <sub>2</sub> O	<50 ppm	100	150	200 `	250	>300 ppm

Milliken & Li, 2017

#### **Exogenous Volatiles**

Solar-wind implanted volatiles (e.g., H, <sup>3</sup>He) in mature, Hi-Ti lunar regolith

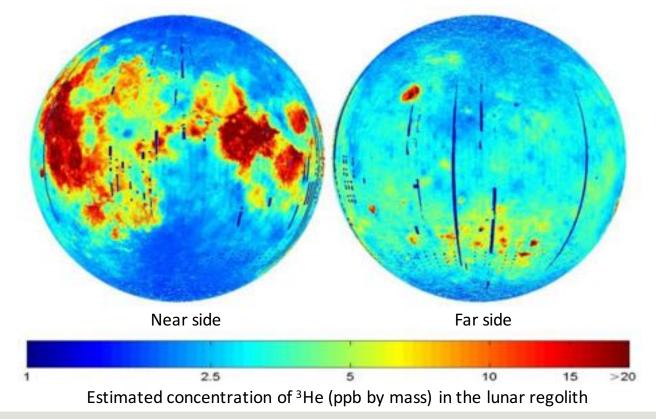
• Evenly distributed in top 2-3 m, requires heating (300°-900° C) to release (Fegley & Swindle,

1993)

Table 1. Average concentrations of solar wind implanted volatiles in the lunar regolith (Fegley and Swindle 1993), where the quoted errors reflect the range (± one standard deviation) of values found at different sampling locations. The corresponding average masses contained within 1 m³ of regolith (assuming a bulk density of 1660 kg m⁻³; Carrier et al., 1991) are also given.

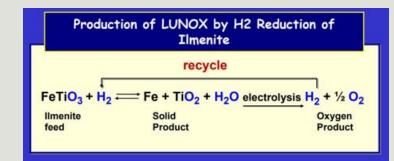
Volatile	Concentration ppm (µg/g)	Average mass per m <sup>3</sup> of regolith (g)
н	46 ± 16	76
<sup>3</sup> He	$0.0042 \pm 0.003$	4 0.007
<sup>4</sup> He	14.0 ± 11.3	23
C	124 ± 45	206
N	81 ± 37	135
F	70 ± 47	116
CI	30 ± 20	50

Crawford, 2015



Fa & Jin, 2007

- Use of in-situ lunar resources can reduce costs of surface operations
  - H<sub>2</sub>O ~ oxygen, drinking water for life support
  - H ~ rocket fuel, reducing agent
  - C, N, S could support lunar agriculture
- Pyroclastic deposits are rich in FeO & TiO<sub>2</sub>
  - Extraction of oxygen (>20 methods to choose from; Taylor & Carrier, 1992)
  - Reduction of ilmenite is popular; requires H, energy
- Survey mode requires mobility, accessibility, traversability
  - Supports horizontal assessment of feedstock, estimates of thickness and vertical distribution from vent outward, determination of consistency of materials
- Sample return provides more precise determinations of composition, calibration for remote mapping
  - Ilmenite content
  - Presence of olivine
  - Indigenous water
  - Surface-correlated and solar-wind-implanted volatiles



From Larry Taylor



3-D Printed Moon Base (ESA)

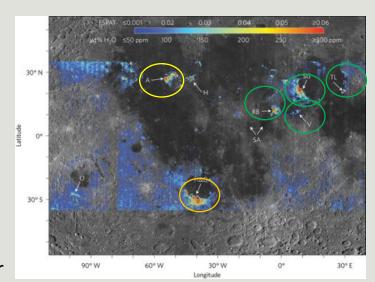
- SKGs to be addressed at lunar pyroclastic deposits
  - Understanding planetary volcanic processes
  - Understanding the Moon's resource potential
  - Understanding the nature & distribution of lunar volatiles

#### Science questions (examples)

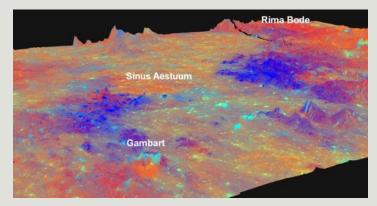
- How old are the lunar pyroclastic deposits?
- Are some eruption styles more likely to be associated with deposits with indigenous water?
- Which deposits contain olivine that can be tied to deep lunar interior origin?
- How thick/uniform are the deposits? Does thickness vary?
   How much material is present?

#### • What sites?

- "Black spot" locations, high iron, titanium, H, H<sub>2</sub>O, surfacecorrelated volatiles, water
  - Taurus-Littrow, Sulpicius Gallus, Rima Bode, Vaporum, etc.
- Humorum
- Aristarchus plateau (Jawin et al., this session)



ESPAT (Milliken & Li, 2017)



3D Sinus Aestuum, Rima Bode (Kaguya MI draped on SLDEM15)

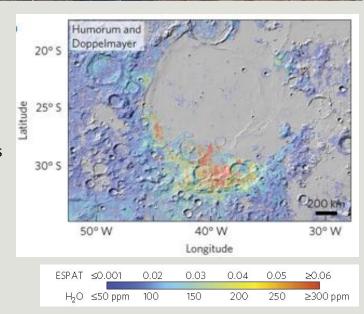
# Thanks for listening!

#### Example Landing Site: Humorum

- Regional extent (1500 km²), smooth-surfaced, Hi-FeO
- Precursor remote characterization
  - High spatial resolution (<100 mpp), hyperspectral imager to map ~3 micron absorption band
    - Characterize abundance, distribution of mafic minerals and indigenous water in detail
  - Use LROC data (~1 mpp) for deposit thickness estimates, boulder& crater hazards
  - Use LOLA data to develop slope maps
- Rover mission (ground contact at multiple sites)
  - Alpha Particle X-ray Spectrometer: Bulk chemistry
  - Neutron Spectrometer: Measure bulk hydrogen, water content at multiple sites
  - Pancam color imager:
    - Multispectral mapping (composition) survey
    - Observe any change(s) with time

#### Sample return

- Multiple samples of mature soil, quenched glass, crystallized beads
- Analysis of magmatic water (abundance, distribution, etc.)
- Ground-truth for remote measurements
- Feedstock assessment and viability for in-situ extraction



Milliken & Li, 2017